



Particulate element composition and phytoplankton biomass in Puddefjorden and Store Lungegårdsvannet during spring 2021

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Abstract

Heavy metal pollution poses a severe threat to aquatic environments, especially in urban coastal areas of high value. This study investigated the elemental composition of particulate matter and the possible impact of heavy metals on phytoplankton biomass in Puddefjorden (PUD) and Store Lungegårdsvannet (SLV), Bergen during spring 2021. Concentrations of C, N, P, Ca, Si, Mg, K, S, Fe, Mn, Cu, Hg, Pb, and As were measured using wavelength dispersive X-ray fluorescence spectroscopy combined with documenting chlorophyll-a concentration. The results revealed that a phytoplankton bloom took place on the 18th of March, coinciding with Si values marking a 42% (PUD) & 50% (SLV) increase, indicating that diatoms dominated the phytoplankton community. Mn, Cu, Hg, Pb, and As were relatively low and did not fluctuate with changes in the phytoplankton biomass, suggesting that these elements did not affect the phytoplankton biomass during the study period. Overall, this study contributes to further understanding of the complex abiotic (particulate chemical elements) – biotic (phytoplankton) relationships in the coastal ecosystem of Bergen.

1. Introduction

Phytoplankton are a driving force of global biogeochemical cycles and play an essential role in the C, P, N, Si cycles in the food webs of aquatic ecosystems (Munn, 2019). Phytoplankton cells are capable of scavenging metal ions from aqueous solutions because of their characteristic rapid growth rate and high surface-to-volume ratio (González-Dávila, 1995; Mehta & Gaur, 2005). They can influence the residual concentration of metal ions by binding them to the cell surface with strong ligands and transporting them intracellularly (Jardim, 1984). Changes in elemental ratios may create unfavorable conditions for phytoplankton growth which in turn may disrupt the marine ecosystem's natural processes. Some ecological effects of the distribution of heavy metals summarized by Brand et al. (1983) include reduced reproductive rates of some phytoplankton species thereby putting a selective pressure on the phytoplankton species populations.

To understand essential ecosystem dynamics like nutrient cycling and food web interactions, it is crucial to broaden our knowledge on how the relationship between abiotic and biotic factors can impact elemental stoichiometry and

vice versa (Erga et al., 2017; Ho et al., 2003; Larsen et al., 2004; Paulino et al., 2013; Price & Skei, 1975). For many decades the Redfield ratio of C:N:P (106:16:1) in phytoplankton biomass and dissolved nutrients has been the holy grail of describing stoichiometric relations of photosynthesis and remineralization (Redfield, 1934). Recent findings, however, show that the ratios of additional chemical elements can reveal information about how phytoplankton biomass is affected by element stoichiometry. The presence of these elements may vary depending on the organisms, the source of detritus or the amount of dissolved organic matter present in the environment, among other things (Arrigo, 2005; Erga et al., 2005; Erga et al., 2017; Erga et al., 2012; Geider & La Roche, 2002; Klausmeier et al., 2004; Larsen et al., 2004; Paulino et al., 2013). For that reason, investigating the elemental composition of suspended particles may unravel a great deal of information about the underlying causes and mechanisms of phytoplankton-elements interactions.

Urban pollution can contribute severely to the input of chemical elements in coastal systems, including essential elements like C, N, and P as well as harmful toxic heavy metals like Hg, Cu, Pb (Andersson & Eggen, 2015; An-

dersson et al., 2014). More information about the availability of elements is needed to understand the impact of their uptake by phytoplankton. In this study, the particulate elemental concentrations of C, N, P, Ca, Si, Mg, K, S, Fe, Mn, Cu, Hg, Pb, and As were determined through wavelength dispersive X-ray fluorescence spectroscopy in combination with documenting chlorophyll-a concentration in Puddefjorden (PUD) and Store Lungegårdsvannet (SLV), located in Bergen. These areas are excellent study subjects as they are a pollution recipient from human activity (Andersson & Eggen, 2015). The findings presented here contributes to the understanding of the complex abiotic (particulate chemical elements) and biotic (phytoplankton) relationships in the coastal ecosystem of Bergen.

2. Study site

PUD is located in the northern part of the Byfjorden water body, which is 3.5 kilometers long, and stretches from the tip of the Nordnes peninsula to the SLV bay (Figure 1). The north-eastern part of PUD is characterized by intense port activity with large commercial vessels, while the inner part is characterized by smaller boats and marinas. Additionally, metal pipelines extend underwater across the seabed. In 2015 a risk assessment of the Puddefjorden seabed concluded that there are concerning levels of mercury pollution associated with the accumulation of pollutants in marine biota. This can pose a severe threat to human health mainly by uptake through seafood consumption (Bergen_Kommune, 2020). In 2018, in the inner part of PUD, a new seabed was overlaid with 45 cm of pure tunnel boring machine masses. Covering the contaminated seabed will, in the long run, act as a physical barrier and

will insulate the underlying contaminated sediment. This action is part of the project *Cleaner Harbor Bergen* which aims to alleviate the urban coastal sites of Bergen Harbor from chronic pollution by organic pollutants and heavy metals so as to limit the content of environmental toxins in fish and seafood from Byfjorden as well as make the area safe for leisure activities (Bergen_Kommune, 2020).

SLV (Figure 1) is also one of the sub-areas in the *Cleaner Harbor Bergen* project, where it is planned to implement measures to monitor the seabed heavy metal pollution in 2022. It has a maximum depth of 26 m and is situated at the southeast end of PUD. The site has a long history as an industrial area, including milling in the Middle Ages, rapid urban and industrial development along the shore, increased shipping activity (Andersson et al., 2014), road traffic, old paint, and other façade materials (Bergen_Kommune, 2020). The water mass of SLV has been dramatically modified over the last 150 years due to continuous land fillings associated with rapid urban development, with the most remarkable one in 1926, when the strait between Lille Lungegårdsvann and SLV was completely filled (Andersson et al., 2014). Today, the connection between SLV and PUD is approximately 3.7 m deep. There is a daily water mass exchange between PUD and SLV during the high and the low tides. As a result of the shallow connection, water exchange is limited, and oxygen in the deepest parts of SLV is depleted (Bergen_Kommune, 2020). The Møllendal River originating from the Svartediket reservoir is SLV's primary freshwater source. The flow is generally low but can increase significantly after heavy rainfall (Paetzel & Schrader, 2003) which enhances the input of natural minerals in the water.

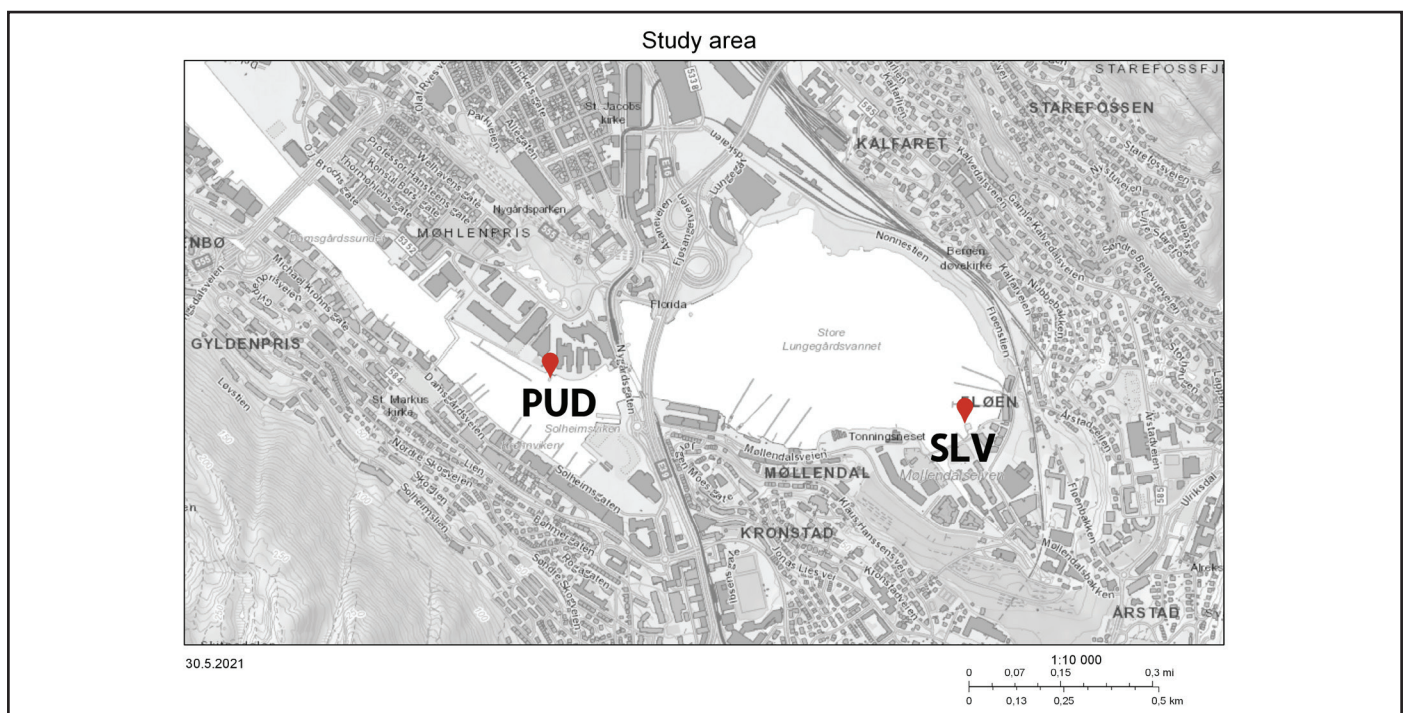


Figure 1. Overview of the study area with the two stations marked on the map, one in front of the Marineholmen Forskningspark in PUD (total depth 5 m) and one in the Møllendal swimming platform in SLV (total depth 11 m).

3. Materials and methods

3.1. Sampling and laboratory analysis

This study was carried out weekly in SLV and PUD from the 26th of February to the 28th of April 2021. Five liters of water samples were collected from each station using a Niskin water sampler at 5 m depth in SLV and 3 m depth in PUD (Figure 1). The water samples were analyzed to measure the chlorophyll a concentration and the particulate elemental composition.

The salinity was recorded using a Xindacheng ATC handheld refractometer. Phytoplankton was identified using a Zeiss Supra 55VP scanning electron microscope after filtering 250 ml of sample material on 1 µm pore size PC filters (25 mm diameter). The chlorophyll a concentration was determined according to Parsons et al. (1984) and Holm-Hansen and Riemann (1978). For each station, three replicates were obtained by filtering 250 ml of sample water onto 0.2 µm Nucleopore polycarbonate filters (PC) (47 mm diameter) using moderate vacuum pressure (≤200 mmHg). The day before the analysis, 10 ml of 100% methanol were added to the samples for chlorophyll-a extraction (for 15-24 hours at ~4°C). The fluorescence was measured with a Turner Designs Fluorometer 10-AU. The whole process took place away from external light sources as far as it was possible. The concentration of chlorophyll-a was calculated according to Equation 1.

Equation 1. Formula for calculation of chlorophyll-a concentration (Holm-Hansen and Riemann 1978; Parsons et al. 1984)

$$Chla = F_d \times \tau \times (R_b - R_a) \times \frac{\text{methanol volume}}{\text{filtration volume}}$$

Where:

Chla: chlorophyll a concentration measured in mg/m³

F_d: acid factor equal to 2.27

$$\tau: \frac{(R_b - R_a)}{(R_b - R_a) - 1} = 1.62$$

R_b: fluorescence value prior to 10% HCl addition

R_a: fluorescence value after 10% HCl addition

methanol volume: 10 ml of 100% methanol

filtration volume: 250 ml from the sample

Water samples for total particulate element analysis were filtered in triplicate (n=3) on 0.6 µm pore size polycarbonate (PC) filters (47 mm diameter), and Whatman glass fiber (GF) filters (57 mm diameter), using ≤200 mmHg vacuum pressure. After filtration, the filters were rinsed with 5 ml of distilled water to prevent interference from salt crystals and left to air-dry until analysis. Total particulate element concentration was measured by wavelength dispersive X-Ray fluorescence spectroscopy (WDXRF) in a Bruker AXE S4 pioneer XRF instrument, based on the method described by Paulino et al. (2013). Data for S, Na, Mg, Cl, K, Ca, Si, Al, Fe, Cu, and Mn were collected from polycarbonate filters, while data on C, N, and

P were taken on three dates (10th of March, 31st of March and 21st of April) from GF filters to avoid interference from the filter's chemical composition. The concentration of all the above-mentioned chemical elements was calculated by Equation 2 and based on the calibration parameters and detection limits explained in Paulino et al. (2013). For Hg, As, and Pb, information about their relative concentrations were recorded manually by looking at the intensity counts (kilo counts per second – kcps) of the K electron shell for Hg and As and the L electron shell for Pb. The criteria chosen for the analysis was the Highest Resolution (HR) which selects the spectral line that favors the best resolution at the expense of intensity. For both methods, at least three blank filters were analyzed for reference. The instrument measures the total amount of an element and does not discern between different chemical forms.

Equation 2. Formula and parameters used for the calculation of the particulate element concentration with WDXRF spectroscopy (Paulino et al. 2013).

$$\text{Concentration} = \left[\frac{(\text{Count}_{\text{sample}} - \text{Count}_{\text{blank}}) \times (\text{filter area}) \times (1000^2 / \text{Volume}) \times (\text{Calibration constant}) \times (\text{Drift ratio})}{\text{Molar Weight}} \right]$$

Where:

Concentration: µM

Count_{sample}: Intensity count (kcps) from the sample

Count_{blank}: Intensity count (kcps) from the blank filter

filter area: 9.078935

Volume: volume of filtered sample

3.2. Statistical analysis

A two-sample t-test was applied to compare the mean concentrations of elements between the two sites. Additionally, the results were analyzed for significant changes in the measurements within each site and between sampling dates, using a two-way analysis of variance (two-way ANOVA). The response variables were the measured concentrations of elements and chlorophyll a, and the predictors were the date and station. Whenever the interaction term was not significant, a backward elimination technique was used to remove those features that did not significantly affect the dependent variable (p > 0.05). All statistical analyses and plotting of the data were performed in RStudio using R 4.1.0 (R Core Team, 2017).

4. Results

Throughout the study period, salinity levels varied between 25-33‰. From the beginning of March, there was a slight differentiation between the two sites, with SLV having a higher average salinity of 30 in 5 m depth, except for the 24th of March, where salinity reached 33‰ in PUD at 3 m depth. During the second half of the study period, salinity in both areas was on average 27‰.

At the beginning of the study, the chlorophyll-a concen-

tration was 1.05 mg/m^3 and 0.51 mg/m^3 in PUD and SLV, respectively. The maximum chlorophyll a concentrations were measured on the 18th of March in both sites (PUD: 16.46 mg/m^3 , SLV: 18.83 mg/m^3), followed by a dramatic drop on the 24th of March (Figure 2). Subsequently, the concentration marked an upward trend until the 21st of April. A two-way ANOVA revealed that the variations in chlorophyll a were statistically significant within each area between each date ($p < 0.05$). A two-sample t-test comparing the total average concentration between the two areas showed no statistically significant difference ($p > 0.05$).

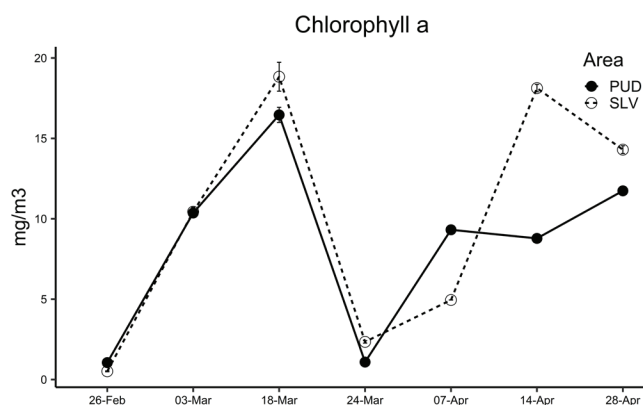


Figure 2. Chlorophyll-a concentration from PUD and SLV during 26th of February – 28th of April 2021.

Particulate C, N, and P average values were $19.76 \text{ }\mu\text{M}$, $2.74 \text{ }\mu\text{M}$, $0.35 \text{ }\mu\text{M}$ in PUD and $22.12 \text{ }\mu\text{M}$, $3.69 \text{ }\mu\text{M}$, $0.43 \text{ }\mu\text{M}$ in SLV, respectively. The lowest values in PUD were recorded on the 10th of March, while for SLV on the 31st of March. Generally, C, N, and P concentrations fluctuated in agreement with each other, with N and P remaining relatively stable. On the 31st of March, maximum C:N ratios were recorded in both areas (PUD: 8.88 and SLV: 7.62),

which was higher than the Redfield ratio (6.6). The maximum C:P ratios were measured on the 21st of April in both sites (PUD: 59.4 and SLV: 68.85). During the study, average values of C:N, C:P, and N:P ratios were 7.34, 7.3, 53.04 in PUD and 6.32, 8.62, 53.15 in SLV, respectively (Figure 3).

Particulate Ca and Si fluctuated between $0.15\text{--}0.95 \text{ }\mu\text{M}$ and $0.96\text{--}4.57 \text{ }\mu\text{M}$, noting a maximum on the 28th of April and 18th of March, respectively. Particulate Ca varied significantly only between the sampling dates ($p < 0.05$), while Si was considerably higher in SLV throughout the study ($p < 0.05$). The maximum particulate Si concentrations observed on the 18th of March coincided with a bloom of the diatom *Skeletonema marinoi*. S concentration was generally higher in SLV ($p = 0.05$) but remained low throughout the study. Fe concentration during the first three weeks and the last two weeks of the investigation varied between $0.19\text{--}0.74 \text{ mM}$ in both locations but marked a maximum of 3.74 on the 31st of March in SLV. Particulate K and Mg remained at low concentrations and peaked during the second half of the study period but differed significantly between the two areas and between each sampling date ($p < 0.05$). Maximum K and Mg values reached $0.47 \text{ }\mu\text{M}$ in SLV on the 28th of April and $0.64 \text{ }\mu\text{M}$ in PUD on the 31st of March, respectively (Figure 4).

The particulate concentration of Cu and Mn $0.002\text{--}0.004 \text{ }\mu\text{M}$ and $0.007\text{--}0.281 \text{ }\mu\text{M}$ (Table 1). However, there was no significant fluctuation found or difference between the areas ($p > 0.05$), suggesting there was little to no variation in the concentration of these metals.

Statistical analysis of the intensity counts for Hg, Pb, and As, revealed no significant fluctuations throughout the study period ($p > 0.05$), although the

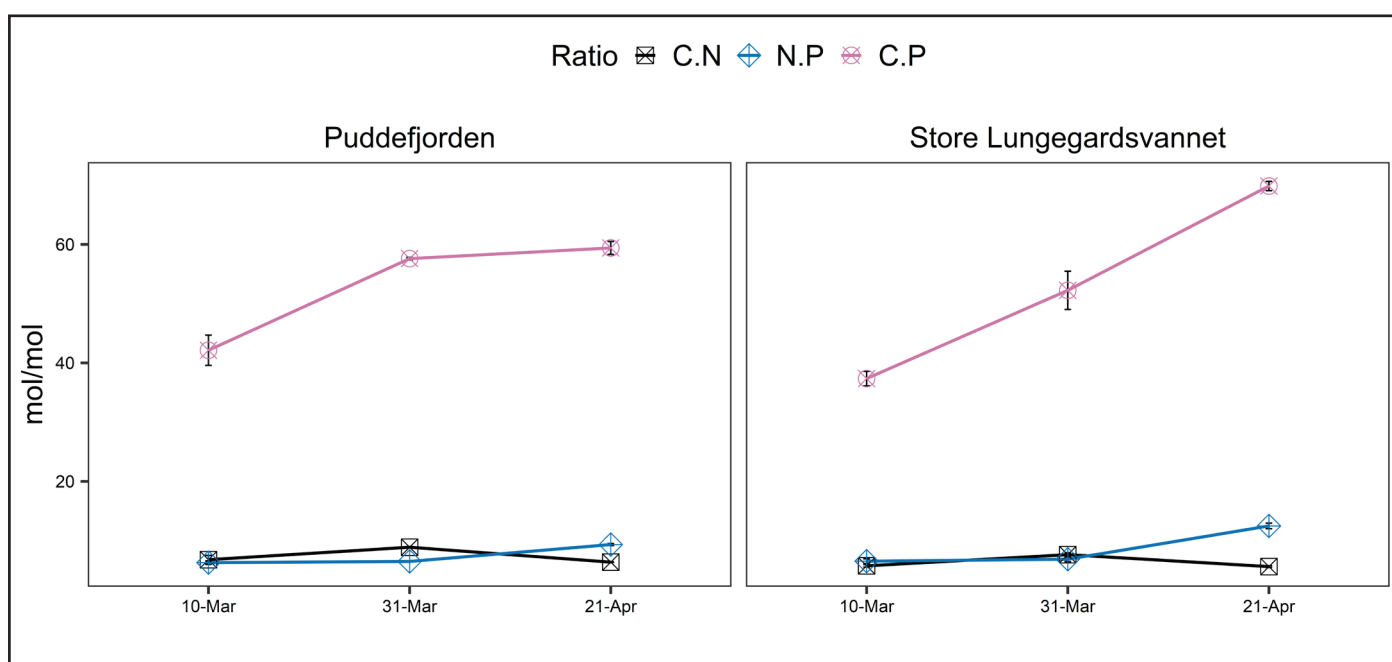


Figure 3. Elemental ratios C:N, N:P and C:P in PUD and SLV during spring 2021.

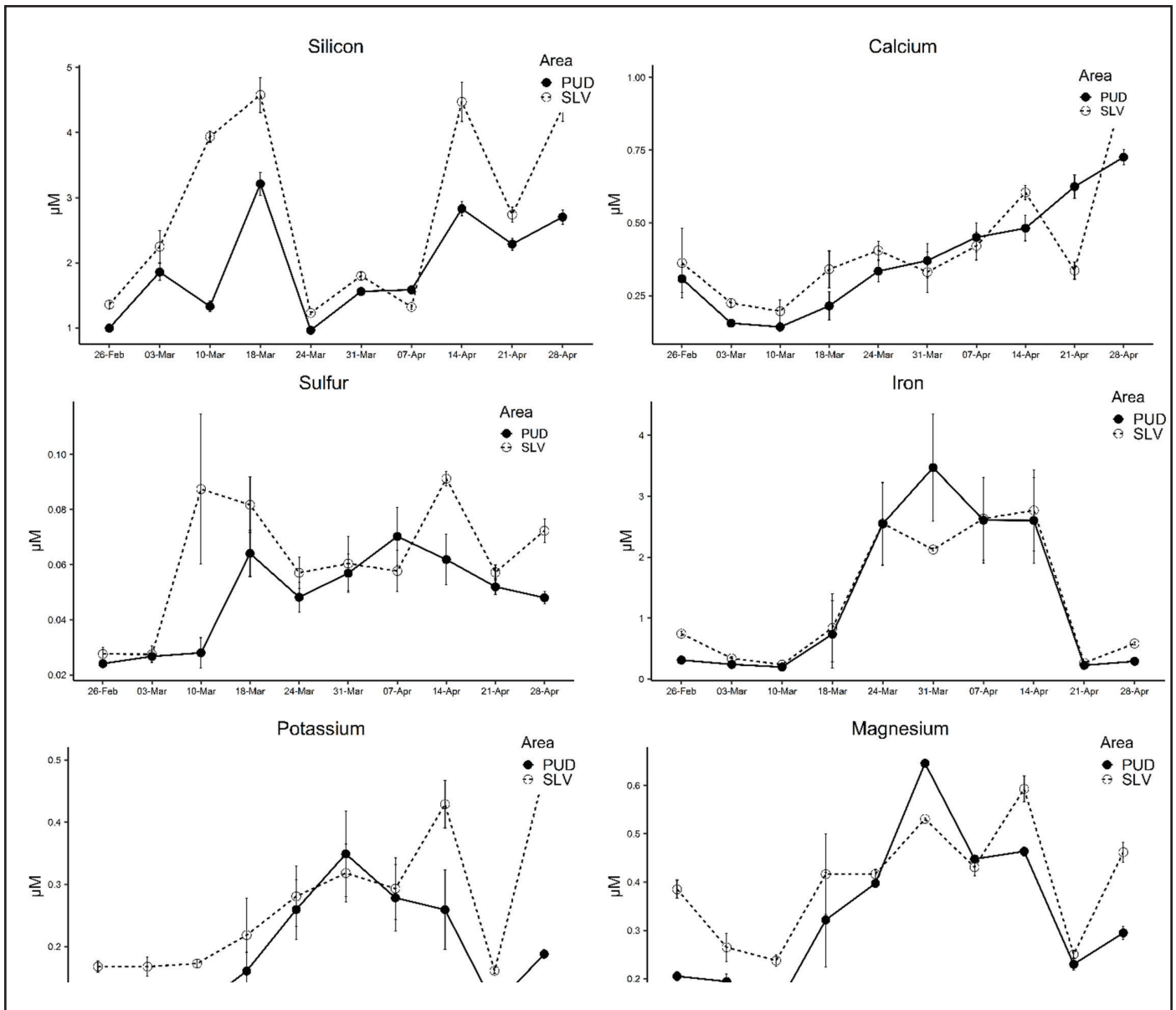


Figure 4. Elemental composition of seston from 26th of February to 28th of April 2021 in PUD and SLV.

counts of all elements were higher in SLV ($p > 0.05$).

5. Discussion

5.1. Low Redfield ratios explained by natural processes and anthropogenic P input

In spring 2021, in both study sites, a diatom bloom took place on the 18th of March, which also occurred during the annual spring diatom bloom in Raunefjorden, which is typically taking place shortly after the 5th of March (Howe et al., 2010; Paulino et al., 2013). Phytoplankton play a major role in the carbon cycle as primary producers (Arrigo, 2005). Other essential micronutrients for cell growth are P and N, which are available through biochemical cycling as well (Arrigo, 2005). Thus, seston concentrations of C, N, and P can be particularly informative about the trophic state of an area when inspected in relation to one another. However, the Redfield atomic ratio interpretation can differ depending on the location, the season,

and the ecological condition (Geider & La Roche, 2002). For example, Norwegian fjords are typically described as P-limited due to the deficit of phosphorus from freshwater sources (Erga et al., 2017; Erga et al., 2012; Paasche & Erga, 1988; Sakshaug et al., 1983; Thingstad et al., 1993). All planktonic organisms contain C, N, and P, and populations of autotrophic and heterotrophic organisms will contribute to these element pools (Erga et al., 2017).

In this study, the C:N:P ratio for PUD was 53:7:1 and 53:9:1 for SLV. C, N, and P peaked around the time of the diatom bloom, suggesting that the contribution by diatoms to the concentrations of these elements was significant. Erga et al. (2017) attained a C:N:P-ratio of 66:11:1 (C:N = 6.2) on a 5 m depth in Raunefjord, but concluded that this evidence did not support any significant nutritional deficiency for N or P. On the contrary, they found that a surplus of P likely was a response to the extreme negative North Atlantic Oscillation (NAO) in 2009-2010. Additionally, they

Date	Cu (Mean \pm SD)		Mn (Mean \pm SD)	
	PUD	SLV	PUD	SLV
26-Feb	0.0036 \pm 0.0003	0.0035 \pm 0	0.0121 \pm 0.0009	0.0126 \pm 0.0016
03-Mar	0.003 \pm 0.001	0.0024 \pm 0.0004	0.0082 \pm 0.0012	0.0119 \pm 0.0015
10-Mar	0.0025 \pm 0.0001	0.0033 \pm 0.0002	0.0079 \pm 0.001	0.0105 \pm 0.0007
21-Apr	0.003 \pm 0.0003	0.003 \pm 0.0007	0.0082 \pm 0	0.0108 \pm 0.0003
28-Apr	0.0026 \pm 0.0006	0.0049 \pm 0.0007	0.0117 \pm 0.0008	0.0281 \pm 0.0014
Average	0.0033 \pm 0.0005	0.0035 \pm 0.0004	0.0094 \pm 0.0008	0.0143 \pm 0.0011

Table 1. Particulate concentration of Mn and Cu in PUD and SLV during spring 2021.

concluded that great values of C: N and C:P atomic ratios were linked to haline stratification and high chlorophyll-a concentration, whereas high N:P ratios are linked to water column mixing with low phytoplankton biomass. Atomic N:P ratios of seston matter that divert from the Redfield ratio could also be attributed to the different species that may dominate the phytoplankton community (Klausmeier et al., 2004). Therefore, C:N:P ratios depend highly on the biological, physical, and chemical processes in a certain area.

Another explanation of the low C:N:P in PUD and SLV recorded in this study could be due to a P surplus. In their research, Martiny et al. (2013) stated that elemental ratios followed a latitudinal pattern due to variations in the coupling of C, N, and P cycles along the different latitudes and observed a C:N:P ratio of 78:13:1 in cold, nutrient-rich high-latitude areas. According to the hypothesis by Redfield (1934), N:P ratio below 16 assumes N-limitation (Frigstad et al., 2011; Geider & La Roche, 2002). However, as Geider & La Roche (2002) suggested, according to studies comparing marine phytoplankton elemental composition, a ratio of N:P<16 of oceanic particulate matter can imply P-rich waters, since it can range from 5 to 19 under sufficient nutrient conditions. However, more evidence is required to determine whether a low C:N:P ratio could be attributed to P-surplus due to anthropogenic nutrient input or other physical-chemical processes in these areas.

5.2. Increased silicate as the aftermath of a *Skeletonema marinoi* bloom

Si and Ca are fundamental components of cellular structures in many plankton taxa like coccolithophorids, foraminifera, and diatoms (Buchan et al., 2014; Howe et al., 2010; Langer, 2008). High concentrations of these elements are linked to blooms of these organisms (Buchan et al., 2014; Howe et al., 2010; Paulino et al., 2013). In April 1973, in Hardangerfjord, particulate Ca and Si values were approximately 0.18 μ M and 4.5 μ M (Erga et al., 2017; Price & Skei, 1975). In 2008, Raunefjorden, particulate Si concentrations corresponded well with the two distinct diatom

blooms, both dominated by *Skeletonema* sp. (Erga et al., 2017). In late spring 2010, at Raunefjorden, Ca and Si varied between 0.099–6.28 μ M and 0.071–3.63 μ M, respectively, coinciding with the high abundance of diatom species like *Chaetoceros* spp., *Skeletonema marinoi*, and *Pseudonitzschia* sp. (Erga et al., 2017). In this study, a strong link between Si concentrations (0.96–4.57 μ M) and *Skeletonema marinoi* dominated bloom was observed. Additionally, the study area is directly influenced by the freshwater input from the Møllendal River on the southeast part of the bay, which is an important source of particulate Si in the fjord (Paetzel & Schrader, 2003). Regarding particulate Ca, biogenic matter is probably the primary source, basically in the form of coccoliths. *Emiliania huxleyi* is one of the most common calcifying species in Norwegian coastal waters, responsible for blooms following the spring diatom bloom (Fagerbakke et al., 1994; Heimdal et al., 1994). However, it seems unlikely that *E. huxleyi* would be entirely responsible for the Ca concentrations (0.15–0.95 μ M) observed in this study but may be also attributed to the water mixing and/or the riverine input. The results from this study for particulate K and Mg agree with results reported in previous studies at nearby fjords (Erga et al., 2017; Price & Skei, 1975).

5.3. Metal concentrations

Fe is thought to limit primary production in marine oceanic systems (Wells et al., 1995). In this study, Fe concentration ranged between 0.19–3.47 μ M, which is higher than the values typically observed in Norwegian fjords. Due to a handling error during the measurement process for the samples from 18th of March to 14th of April, the concentration for elements with high molar weight like Fe, Mn, Cu do not reflect their actual particulate concentration, and are thus not reliable. That is the primary reason why maximum Fe values documented here greatly exceed the typical particulate Fe concentrations in marine seston. Typical values for fjords on the west coast of Norway range between 0.068–0.23 μ M (Price & Skei, 1975), however, in areas strongly influenced by terrigenous sources, the concentration of particulate Fe can vary between 0.04–2.8 μ M

(Bazzano et al., 2014). In a study performed on shelf waters of the eastern United States, concentrations of particulate Fe and Al were indirectly linked to salinity since their concentrations in runoff water can be several orders of magnitude higher than in oceanic waters (Atkinson & Stefánsson, 1969). Erga et al. (2017) concluded that the prominent iron sources in the samples from Raunefjorden 2009-2011 were Fe sequestering bacteria (Heldal et al., 1996), Fe bound to phytoplankton cells, as well as Fe fractions associated with certain microalgae blooms. All these factors could explain the Fe values observed in this study, together with some other exogenous sources from contamination in the study area; however, this is subject to further research.

So far, no comprehensive studies have been performed on the particulate concentration of heavy metals in the study area. Most of the studies focus on analyzing the sediment and providing estimates of the pollution state based on these values. Comparison of results from sediment surveys in the entire PUD between 1992 and 2020 (Table 2) shows a remarkable decrease in the sediment concentration of most heavy metals in PUD, especially after covering the seabed in 2018 (Bergen_Kommune, 2020). In this study, I showed that heavy metal presence (Hg, Pb, As) was significantly different in the two sites ($p < 0.05$) with PUD having the lowest amount. It is highly possible that the low particulate concentration documented in this study is linked to their concentration in the sediment. A recent follow-up survey in PUD, however, revealed that despite the low level of pollution achieved, the top layer of the seabed is subject to recontamination with environmental toxins from the surrounding urban environment (Bergen_Kommune, 2020). This highlights that proper control of pollution sources that can lead to recontamination is the prerequisite for a monitoring plan to be successful.

6. Conclusions

This study was an effort to document the particulate concentration of chemical elements and associate these concentrations with the phytoplankton biomass. Chlorophyll-a varied with the concentration of C, N, P, and

Si. These elements are important for the development of phytoplankton, especially Si for the development of diatoms that dominated during the spring bloom. Heavy metals were detected in both areas in low levels; however, the concentrations were lower in PUD, suggesting that the municipality's plan to cover the polluted sediment may have resulted in fewer particles in the water column containing those metals. Overall, this study was an example of investigating abiotic and biotic relationships in an urban coastal environment, however, the collection of more observational data is necessary to understand the aftermath of heavy metal pollution and to design effective and sustainable management plans in the long term.

7. References

- Andersson, M., & Eggen, O. A. (2015). Urban contamination sources reflected in inorganic pollution in urban lake deposits, Bergen, Norway. *Environmental Science: Processes & Impacts*, 17(4), 854-867. <https://doi.org/10.1039/c4em00614c>
- Andersson, M., Klug, M., Eggen, O. A., & Ottesen, R. T. (2014). Polycyclic aromatic hydrocarbons (PAHs) in sediments from lake Lille Lungegårdsvannet in Bergen, western Norway; appraising pollution sources from the urban history. *Science of The Total Environment*, 470-471, 1160-1172. <https://doi.org/10.1016/j.scitotenv.2013.10.086>
- Arrigo, K. R. (2005). Marine microorganisms and global nutrient cycles. *Nature*, 437(7057), 349-355. <https://doi.org/10.1038/nature04159>
- Atkinson, L. P., & Stefánsson, U. (1969). Particulate aluminum and iron in sea water off the southeastern coast of the United States. *Geochimica et Cosmochimica Acta*, 33(11), 1449-1453. [https://doi.org/10.1016/0016-7037\(69\)90184-7](https://doi.org/10.1016/0016-7037(69)90184-7)
- Bazzano, A., Rivarolo, P., Soggia, F., Ardini, F., & Grotti, M. (2014). Anthropogenic and natural sources of particulate trace elements in the coastal marine environment of Kongsfjorden, Svalbard. *Marine Chemistry*, 163, 28-35. <https://doi.org/10.1016/j.marchem.2014.04.001>
- Bergen_Kommune. (2020). Renere havn Bergen.
- Brand, L. E., Sunda, W. G., & Guillard, R. R. L. (1983). Limitation of marine phytoplankton reproductive rates by zinc, manganese, and iron1. *Limnology and Oceanography*, 28(6), 1182-1198. <https://doi.org/10.4319/lo.1983.28.6.1182>
- Buchan, A., Leclair, G. R., Gulvik, C. A., & González, J. M. (2014). Master recyclers: features and functions of bacteria associated with phytoplankton blooms. *Nature Reviews Microbiology*, 12(10), 686-698. <https://doi.org/10.1038/nrmicro3326>

Chemical element	1992-2003 (mg/kg)	2005-2014 (mg/kg)	2019-2020 (mg/kg)
As	21	18	1.3
Pb	298	242	16
Cu	414	299	13
Hg	6.2	3.3	0.05
Zn	734	479	38

Table 2. Sediment data used in the risk assessment for PUD divided into three periods. Data from 1992-2003 and 2005-2014 illustrate the condition in the whole fjord, and it is uncertain on which depth and exact sites they were collected. Data from 2019-2020 were taken from the top 10 cm of the seabed in the most inner part of PUD(Bergen Kommune 2020).

- Clarence, R. A. (1934). On the proportions of organic derivatives in sea water and their relation to the composition of plankton. James Johnstone memorial volume, 16.
- Erga, Sr., Aursland, K., Frette, Ø., Hamre, B., Lotsberg, J., Stamnes, J., . . . Stamnes, K. (2005). UV transmission in Norwegian marine waters: controlling factors and possible effects on primary production and vertical distribution of phytoplankton. *Marine Ecology Progress Series*, 305, 79-100. <https://doi.org/10.3354/meps305079>
- Erga, S. R., Haugen, S. B., Bratbak, G., Egge, J. K., Heldal, M., Mork, K. A., & Norland, S. (2017). Seasonal variations in C:N:Si:Ca:P:Mg:S:K:Fe relationships of seston from Norwegian coastal water: Impact of extreme offshore forcing during winter-spring 2010. *Marine Chemistry*, 196, 1-12. <https://doi.org/10.1016/j.marchem.2017.07.001>
- Erga, S. R., Ssebiyonga, N., Frette, Ø., Hamre, B., Aure, J., Strand, Ø., & Strohmeier, T. (2012). Dynamics of phytoplankton distribution and photosynthetic capacity in a western Norwegian fjord during coastal upwelling: Effects on optical properties. *Estuarine, Coastal and Shelf Science*, 97, 91-103. <https://doi.org/10.1016/j.ecss.2011.11.034>
- Fagerbakke, K. M., Heldal, M., Norland, S., Heimdal, B. R., & Båtvik, H. (1994). *Emiliania huxleyi*. Chemical composition and size of coccoliths from enclosure experiments and a Norwegian fjord. *Sarsia*, 79(4), 349-355. <https://doi.org/10.1080/00364827.1994.10413566>
- Frigstad, H., Andersen, T., Hessen, D. O., Naustvoll, L. J., Johnsen, T. M., & Bellerby, R. G. J. (2011). Seasonal variation in marine C:N:P stoichiometry: can the composition of seston explain stable Redfield ratios? *Biogeosciences*, 8(10), 2917-2933. <https://doi.org/10.5194/bg-8-2917-2011>
- Geider, R., & La Roche, J. (2002). Redfield revisited: variability of C:N:P in marine microalgae and its biochemical basis. *European Journal of Phycology*, 37(1), 1-17. <https://doi.org/10.1017/S0967026201003456>
- González-Dávila, M. (1995). The role of phytoplankton cells on the control of heavy metal concentration in seawater. *Marine Chemistry*, 48(3-4), 215-236. [https://doi.org/10.1016/0304-4203\(94\)00045-f](https://doi.org/10.1016/0304-4203(94)00045-f)
- Heimdal, B. R., Egge, J. K., Veldhuis, M. J. W., & Westbroek, P. (1994). The 1992 Norwegian *Emiliania huxleyi* experiment. An overview. *Sarsia*, 79(4), 285-290. <https://doi.org/10.1080/00364827.1994.10413560>
- Heldal, M., Norland, S., Fagerbakke, K. M., Thingstad, F., & Bratbak, G. (1996). The elemental composition of bacteria: A signature of growth conditions? *Marine Pollution Bulletin*, 33(1-6), 3-9. [https://doi.org/10.1016/s0025-326x\(97\)00007-6](https://doi.org/10.1016/s0025-326x(97)00007-6)
- Ho, T.-Y., Quigg, A., Finkel, Z. V., Milligan, A. J., Wyman, K., Falkowski, P. G., & Morel, F. M. M. (2003). THE ELEMENTAL COMPOSITION OF SOME MARINE PHYTOPLANKTON1. *Journal of Phycology*, 39(6), 1145-1159. <https://doi.org/10.1111/j.0022-3646.2003.03-090.x>
- Holm-Hansen, O., & Riemann, B. (1978). Chlorophyll a determination: improvements in methodology. *Oikos*, 438-447.
- Howe, J. A., Austin, W. E. N., Forwick, M., Paetzel, M., Harland, R., & Cage, A. G. (2010). *Fjord systems and archives: a review*. Geological Society, London, Special Publications, 344(1), 5-15. <https://doi.org/10.1144/sp344.2>
- Jardim, W. (1984). A study of the copper-complexing compounds released by some species of cyanobacteria. *Water Research*, 18(8), 985-989. [https://doi.org/10.1016/0043-1354\(84\)90249-5](https://doi.org/10.1016/0043-1354(84)90249-5)
- Klausmeier, C. A., Litchman, E., Daufresne, T., & Levin, S. A. (2004). Optimal nitrogen-to-phosphorus stoichiometry of phytoplankton. *Nature*, 429(6988), 171-174. <https://doi.org/10.1038/nature02454>
- Langer, M. R. (2008). Assessing the Contribution of Foraminiferan Protists to Global Ocean Carbonate Production. *Journal of Eukaryotic Microbiology*, 55(3), 163-169. <https://doi.org/10.1111/j.1550-7408.2008.00321.x>
- Larsen, A., Flaten, G. A. F., Sandaa, R.-A., Castberg, T., Thyraug, R., Erga, S. R., . . . Bratbak, G. (2004). Spring phytoplankton bloom dynamics in Norwegian coastal waters: Microbial community succession and diversity. *Limnology and Oceanography*, 49(1), 180-190. <https://doi.org/10.4319/lo.2004.49.1.0180>
- Martiny, A. C., Pham, C. T. A., Primeau, F. W., Vrugt, J. A., Moore, J. K., Levin, S. A., & Lomas, M. W. (2013). Strong latitudinal patterns in the elemental ratios of marine plankton and organic matter. *Nature Geoscience*, 6(4), 279-283. <https://doi.org/10.1038/ngeo1757>
- Mehta, S. K., & Gaur, J. P. (2005). Use of Algae for Removing Heavy Metal Ions From Wastewater: Progress and Prospects. *Critical Reviews in Biotechnology*, 25(3), 113-152. <https://doi.org/10.1080/07388550500248571>
- Munn, C. B. (2019). *Marine microbiology: ecology & applications*. CRC Press.
- Paasche, E., & Erga, S. R. (1988). Phosphorus and nitrogen limitation of phytoplankton in the inner Oslofjord (Norway). *Sarsia*, 73(3), 229-243. <https://doi.org/10.1080/00364827.1988.10413409>
- Paetzel, M., & Schrader, H. (2003). Natural vs. human-induced facies changes in recent, shallow fjord sediments of the Store Lungegårdsvannet in Bergen (western Norway). *Environmental Geology*, 43(4), 484-492. <https://doi.org/10.1007/s00254-002-0663-3>
- Parsons, T. R., Maita, Y., & Lalli, C. M. (1984). *A manual of chemical and biological methods for seawater analysis* (1st ed.). Pergamon Press.
- Paulino, A. I., Heldal, M., Norland, S., & Egge, J. K. (2013). Elemental stoichiometry of marine particulate matter measured by wavelength dispersive X-ray fluorescence (WDXRF) spectroscopy. *Journal of the Marine Biological Association of the United Kingdom*, 93(8), 2003-2014. <https://doi.org/10.1017/s0025315413000635>
- Price, N. B., & Skei, J. M. (1975). Areal and seasonal variations in the chemistry of suspended particulate matter in a deep water fjord. *Estuarine and Coastal Marine Science*, 3(3), 349-369. [https://doi.org/10.1016/0302-3524\(75\)90034-1](https://doi.org/10.1016/0302-3524(75)90034-1)
- R Core Team. (2017). R: A language and environment for statistical computing. In R Foundation for Statistical Computing. <https://www.R-project.org/>
- Redfield, A. C. (1934). On the proportions of organic derivatives in sea water and their relation to the composition of plankton. James Johnstone memorial volume, 16.
- Sakshaug, E., Andresen, K., Mykkestad, S., & Olsen, Y. (1983). Nutrient status of phytoplankton communities in Norwegian waters (marine, brackish, and fresh) as revealed by their chemical composition. *Journal of Plankton Research*, 5(2), 175-196. <https://doi.org/10.1093/plankt/5.2.175>
- Thingstad, T. F., Skjoldal, E. F., & Bohne, R. A. (1993). Phosphorus cycling and algal-bacterial competition in Sandsfjord, western Norway. *MARINE ECOLOGY PROGRESS SERIES*, 99, 20.
- Wells, M. L., Price, N. M., & Bruland, K. W. (1995). Iron chemistry in seawater and its relationship to phytoplankton: a workshop report. *Marine Chemistry*, 48(2), 157-182. [https://doi.org/10.1016/0304-4203\(94\)00055-i](https://doi.org/10.1016/0304-4203(94)00055-i)

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